

Mapping global surface roughness using AMSR-E passive microwave remote sensing



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ARTICLE INFO

Article history:

Received 13 March 2014

Received in revised form 24 July 2014

Accepted 28 July 2014

Available online 8 August 2014

Keywords:

Surface roughness

Physical-based retrieval model

Passive microwave remote sensing

Advanced Microwave Scanning Radiometer –

Earth Observing System (AMSR-E)

Brightness temperature (T_b)

ABSTRACT

Surface roughness is an important geography parameter required in various research fields. Passive microwave remote sensing could be stably used for global surface roughness simulations. This paper develops a physical-based surface roughness retrieval model on the basis of the passive microwave radiative transfer equation using the C-band (6.9 GHz) and X-band (10.7 GHz) brightness temperatures of Advanced Microwave Scanning Radiometer – Earth Observing System (AMSR-E). Results show that there is a significant linear relationship between the model-derived surface roughness and that simulated by Hong (2010a) ($R^2 = 0.62$; root mean square error (RMSE) = 0.09 cm). Seasonal variations of the monthly average surface roughness in 2009 indicate that vegetated lands and freezing lands are usually significantly rougher than the un-vegetated bare grounds, while snow-cover lands and deserts always have smoother surfaces than the vegetated lands. In summary, vegetation cover, freezing soil and snow cover are important factors influencing the land surface roughness conditions. It is worthy mentioning that the proposed global surface roughness retrieval model only uses AMSR-E C- and X-band brightness temperatures without any ancillary data and might be fitted for global applications.

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1. Introduction

Surface roughness is a vital parameter in the pedology, agronomy, geology and hydrology research fields (Singh et al., 2000, 2003). Generally, it has two definitions. One is defined as the large-scale surface roughness that represents the surface topographic variations at length scales below the resolution of available digital elevation models or maps (Mushkin and Gillespie, 2005). Another is defined as the small-scale surface roughness, which represents the standard deviation of the surface heights that caused by soil/rock variations but not including the vegetations (Hong, 2010a). This paper refers to the latter definition. The small-scale surface roughness not only has a significant impact on infiltration and soil erosion in the agronomy and pedology processes, but also influences the ability of soil to store and release water through evaporation in hydrologic processes. In addition, Mama et al. (1997) demonstrated that surface roughness could provide useful information on explaining or distinguishing the geologic process such as eolian sand and rocky surface. Some application studies also found that remote sensing accuracy of surface soil moisture is very sensitive to surface

roughness (Dobson et al., 1985; Fung, 1994; Ulaby and Bare, 1979; Ulaby et al., 1986), which poses major problems for current soil moisture retrievals (Davidson et al., 2000; Mattia and Le Toan, 1999; Wagner et al., 2007). Therefore, it is necessary to accurately map the regional surface roughness conditions (Moran et al., 2004).

During the past 20 years, several important studies have been conducted and results have proved that passive microwave remote sensing could be stably used for global surface roughness simulations. Huang and Jin (1995) firstly employed the scattering theory of rough surface and reciprocity using both the active and passive microwave data. Then, the relationship between surface emissivity and backscattering coefficients for surfaces of various roughness degrees was derived based on the proposed scattering theory. Finally, the surface roughness variables were retrieved from the microwave emissivity using a mesh graph, which was constructed on the basis of the relationship between surface emissivity and backscattering coefficients. But the mesh graph method was a little complicated for global applications. Hong (2010a, 2010b) established an approximate relationship between the vertically and horizontally polarized reflectivities and developed a unique retrieval small-scale roughness model. Then, the global small-scale roughness over land surfaces was estimated using the 6.9 GHz brightness temperature (T_b) of the Advanced Microwave Scanning Radiometer – Earth Observing System (AMSR-E) based on the proposed retrieval model. The small-scale roughness was retrieved within the reasonable range of previous works.

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Hong (2010b) expanded the model and applied for retrieving the small-scale roughness over sea ices using AMSR-E T_b . Results also showed reasonable agreement with the known observations, ranging from 0.2 cm to 0.6 cm for the sea ice in the Antarctic and Arctic regions. However, Hong's method had to use other parameters such as land surface temperature as the input variable for supporting the global surface roughness simulation (Hong, 2010a). Chen et al. (2012) developed a surface roughness index using AMSR-E C-band (6.9 GHz), X-band (10.7 GHz) and L-band (18.7 GHz) Microwave Polarization Difference Index (MPDI). Results showed that the proposed surface roughness index minimized the impacts from the vegetations and was effective to map the surface roughness at both global scale and regional scale. But the empirical surface roughness index also has its defects and might bring some disadvantages when applied for time-series studies at the global scale. In summary, remote sensing of global surface roughness is still rarely studied at present. In particular, most current methods have some disadvantages: either some need variables as the input parameters or some are empirical models, which might bring unexpected errors when applied for long-time applications at the global scale. Hence, new surface roughness retrieval models are needed at the current stage.

In this paper, we aim to develop a physical-based model to map global surface roughness from AMSR-E C- and X-band T_b without any ancillary data. The proposed model is based on the passive microwave radiative transfer equation (Mo et al., 1982; Owe et al., 2001), surface roughness model (Wang and Choudhury, 1995), vegetation optical depth (τ_c) model (De Jeu, 2003) and MPDI (Becker and Choudhury, 1988; Choudhury and Tucker, 1987; Choudhury et al., 1987).

2. Materials

AMSR-E radiometer onboard the NASA Earth Observing System (EOS) Aqua satellite is a modified version of the AMSR radiometer that launched on the Japanese Advanced Earth Observing Satellite-II (ADEOS-II) in 1999. It is a conically scanning total power passive microwave radiometer sensing microwave T_b at 12 channels and 6 frequencies (6.9, 10.65, 18.7, 23.8, 36.5 and 89.0 GHz). All channels operate in both vertical (V) and horizontal (H) polarization at the incidence angle of 55.0°. The National Snow and Ice Data Center (NSIDC) has provided AMSR-E/Aqua Daily Global Quarter-Degree Gridded Brightness Temperatures data in the grid-format file, which are produced in the Equal-Area Scalable Earth Grid (EASE-Grid) at 25 km horizontal resolution in one global cylindrical, equidistant latitude–longitude projection (<http://nsidc.org/>). On October 4th, 2011, the AMSR-E radiometer reached its limit to maintain the rotation speed necessary for regular observations and automatically halted its observations and rotations. The Advanced Microwave Scanning Radiometer 2 (AMSR2) onboard the GCOM-W1 satellite has been successfully spun up, which will provide us with highly accurate measurements of the intensity of microwave emission and scattering. In the study, one-year global AMSR-E T_b data in 2009 downloaded from NSIDC are used to establish the surface roughness retrieval model.

3. Methods

3.1. Basic theory of the small-scale surface roughness model

The small-scale roughness of a material surface can be defined as W-C model (Wang and Choudhury, 1995; Wigneron et al., 2001).

$$\sigma = \frac{\lambda}{4\pi \cos\theta} \sqrt{h} \quad (1)$$

where σ is the small-scale surface roughness, unit: cm; λ is the wavelength; π is the circumference ratio; θ is the incident angle of the remote sensed instrument; h represents the effective roughness height.

The reflectivity of rough surfaces can be expressed by Wang and Choudhury's Q/H surface roughness model (Wang and Choudhury, 1995).

$$r_{sv} = [(1-Q)r_{ov} + Qr_{oh}]e^{-h} \quad (2-1)$$

$$r_{sh} = [(1-Q)r_{oh} + Qr_{ov}]e^{-h} \quad (2-2)$$

where r_{sv} and r_{sh} represent the vertical polarization and horizontal polarization reflectivities of rough surfaces, respectively; r_{ov} and r_{oh} represent the vertical polarization and horizontal polarization reflectivities of flat surfaces, respectively; Q is the cross polarization ratio of surface roughness; h represents the effective roughness height. The values of Q for 6.6 GHz and 10.7 GHz frequencies simulated to equaling 0.09 and 0.11 by Njoku and Li (1999).

3.2. Simplification of the passive microwave radiative transfer equation

A simple atmospheric correction was firstly implemented to eliminate the atmospheric influences on AMSR-E C-band and X-band T_b using Pellarin's atmospheric correction method (Pellarin et al., 2003, 2006), which quantified globally at L-band, C-band and X-band based on the Earth Vegetation Atmosphere Model (EVA, Kerr and Njoku, 1990). The atmospheric effects for the three wavelengths are presented in Fig. 1 and Table 1. In this paper, values of 2.17 and 3.45 K are used to correct the atmosphere effects on AMSR-E C-band and X-band microwave radiations, respectively.

After the simple atmospheric correction, the atmospheric impacts on microwave emission can be ignored in the radiative transfer equation, and the radiative T_b of AMSR-E C-band and X-band is expressed as (Mo et al., 1982):

$$T_{bp} = T_s(1-r_{sp})e^{-\tau_c} + T_c(1-w_p)(1-e^{-\tau_c}) + r_{sp}T_c(1-w_p) \times (1-e^{-\tau_c})e^{-\tau_c} \quad (3)$$

where p is the horizontal (H) or vertical (V) polarization mode; T_{bp} is the brightness temperature at the p polarization mode; T_s represents the thermometric temperatures of the soil; r_{sp} is the rough-surface reflectivity; τ_c is the optical depth of land surface vegetation; $e^{-\tau_c}$ is the transmissivity; T_c is thermometric temperatures of the canopy; w_p is the single scattering albedo.

Many studies have demonstrated that the effects of the single scattering albedo w_p are negligible and can be ignored at low frequencies (<10 GHz, like AMSR-E C- and X-bands) (Chen et al., 2011, 2012;

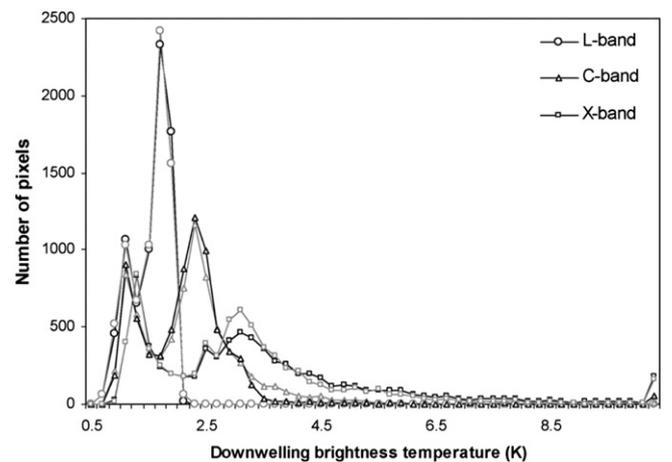


Fig. 1. Global-scale atmospheric influences on the down-welling brightness temperatures (T_b) of microwave L, C and X-band wavelengths, respectively (Pellarin et al., 2006).

Table 1
Average atmospheric contributions on C-band and X-band T_b of 1.4, 6.9 and 10.7 GHz globally (Pellarin et al., 2006).

Microwave frequency (GHz)	Average atmospheric contributions on T_b (K)
1.4	1.44
6.6	2.17
10.7	2.35

Hong and Shin, 2011; Schmugge and Jackson, 1993; Wang et al., 2006). In order to derive a simple estimation model from the radiative transfer equation, we ignore the single scattering contributions on AMSR-E C-frequency (6.9 GHz) and X-frequency (10.7 GHz) and assume that the surface soil temperature T_s equals to the vegetation canopy temperature T_c (Njoku and Li, 1999; Paloscia and Pampaloni, 1988; Schmugge and Jackson, 1993; Wigneron et al., 2003; Meesters et al., 2005; Hong and Shin, 2011). Then, the passive microwave radiative transfer equation is simplified as:

$$T_{bp} = T_c(1 - r_{sp}e^{-2\tau_c}). \tag{4}$$

3.3. Development of the surface roughness retrieval model

MPDI is a classic microwave vegetation index for characterizing the land surface vegetation cover conditions (Jackson and Schmugge, 1991; Kerr and Njoku, 1990; Le Vine and Karam, 1996; Njoku and Li, 1999; Pampaloni and Paloscia, 1986). Its definition formula is described as:

$$MPDI = \frac{T_{bv} - T_{bh}}{T_{bv} + T_{bh}} \tag{5}$$

where T_{bh} and T_{bv} represent the horizontal and vertical T_b , respectively.

Substitute Eqs. (4) and (5) into Eq. (1), leading to (Wang et al., 2006):

$$\frac{1}{MPDI} = \frac{r_{ov} + r_{oh}}{(1-2Q)(r_{ov} - r_{oh})} - \frac{2}{(1-2Q)(r_{ov} - r_{oh})} e^{2\tau_c + h}. \tag{6}$$

De Jeu (2003) developed a relation (Eq. (7)) between the vegetation optical depth (τ_c) and MPDI on the basis Meesters' τ_c model (Meesters

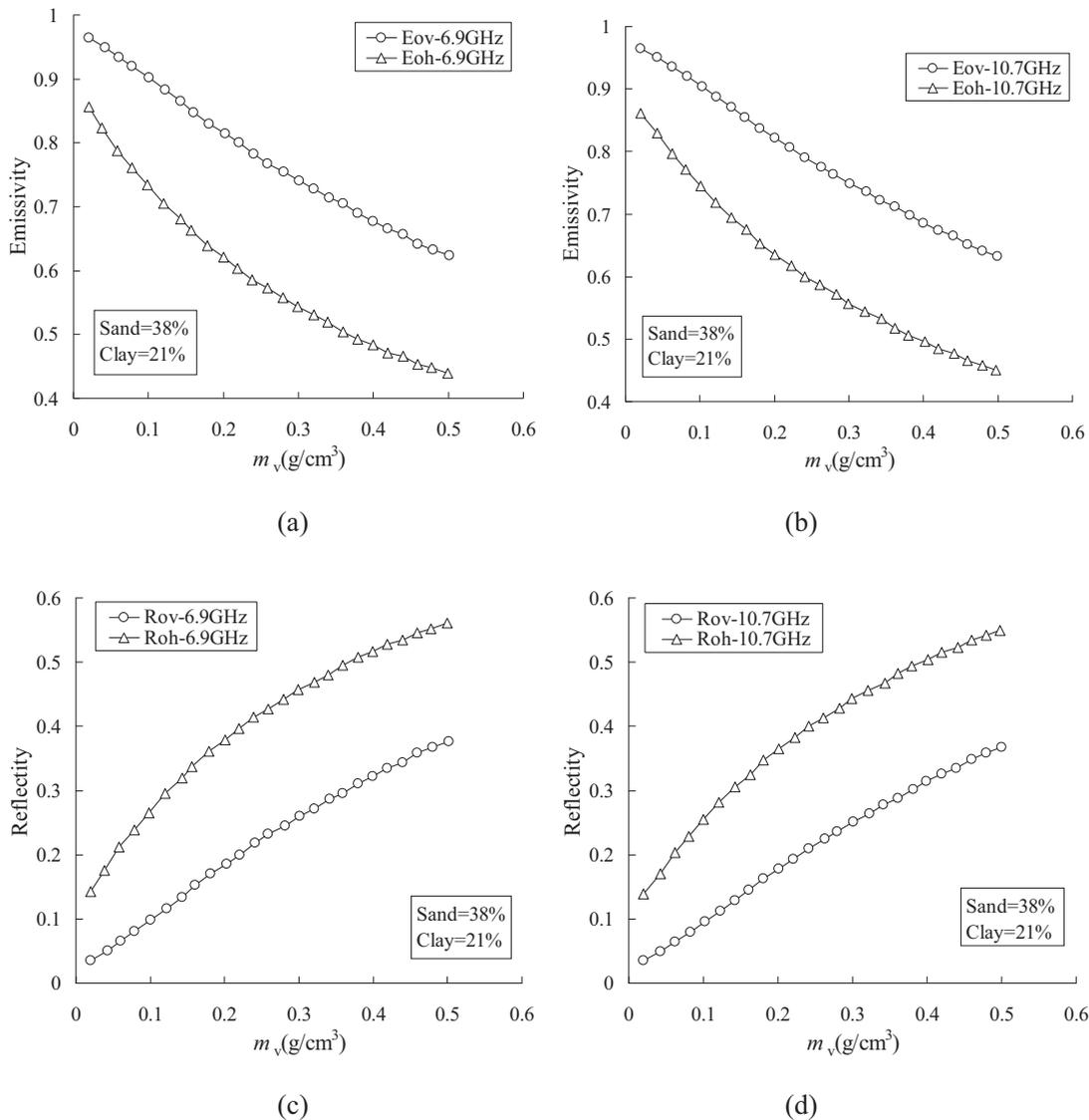


Fig. 2. Relations between surface soil moisture (m_v) and soil emissivity (E_{op}), and m_v and surface reflectivity (r_{op}) for AMSR-E 6.9 and 10.7 GHz vertical (V) and horizontal (H) polarization channels.

Table 2
Values of the parameters in the surface roughness retrieve model (Eqs. (10-1), (10-2)).

Parameters	Values
Q	0.09
α	1.2446
β	0.3586

et al., 2005). Wang et al. (2006) and Ma (2007) have proved that De Jeu's empirical solution of τ_c held well with Wang and Choudhury's Q/H surface roughness model (Eq. (2)).

$$\tau_c = C_1 \ln(\text{MPDI})^3 + C_2 \ln(\text{MPDI})^2 + C_3 \ln(\text{MPDI}) + C_4 \quad (7)$$

where C_1, C_2, C_3 and C_4 are undetermined coefficients.

By substituting Eq. (7) into Eq. (6), we come to:

$$(\text{MPDI} - 1 + 2Q) \times r_{ov} + (\text{MPDI} + 1 - 2Q) \times r_{oh} = 2(\text{MPDI})^\alpha \times e^{\beta+h} \quad (8)$$

where $\alpha = 6C_1 + 4C_2 + 2C_3, \beta = 2C_4$. The optimal solutions for α and β were set as 1.2446 and 0.3586 that were simulated from 1152 pairs of MPDI and surface soil moisture (m_v) values using the Levenberg-Marquardt method (Chen et al., in press).

Further, the vertical polarization emissivity (E_{ov}) and horizontal polarization emissivity (E_{oh}) of AMSR-E C-band (6.9 GHz) and X-band (10.7 GHz) under different m_v from 0 g/cm³ to 0.50 g/cm³ were simulated using the method of Zhong (2005), shown in Fig. 2a and b. Then, the relations between m_v and r_{ov}, m_v and r_{oh} (Fig. 2c and d) are derived from the surface emissivity (E_{op}) using the equation: $r_{op} = 1 - E_{op}$. Relation equations for AMSR-E C- and X-bands are listed as follows:

$$r_{ov6.9} = 0.7258 * m_v + 0.0314 (R^2 = 0.993) \quad (9-1)$$

$$r_{oh6.9} = 0.7757 m_v^{0.4481} (R^2 = 0.997) \quad (9-2)$$

$$r_{ov10.7} = 0.7117 m_v + 0.0284 (R^2 = 0.995) \quad (9-3)$$

$$r_{oh10.7} = 0.7619 m_v^{0.4610} (R^2 = 0.994) \quad (9-4)$$

where $r_{ov6.9}$ means the flat-surface reflectivity of C-band V polarization; $r_{oh6.9}$ means the flat-surface reflectivity of C-band H polarization; $r_{ov10.7}$ means the flat-surface reflectivity of X-band V polarization; and $r_{oh10.7}$ means the flat-surface reflectivity of X-band H polarization.

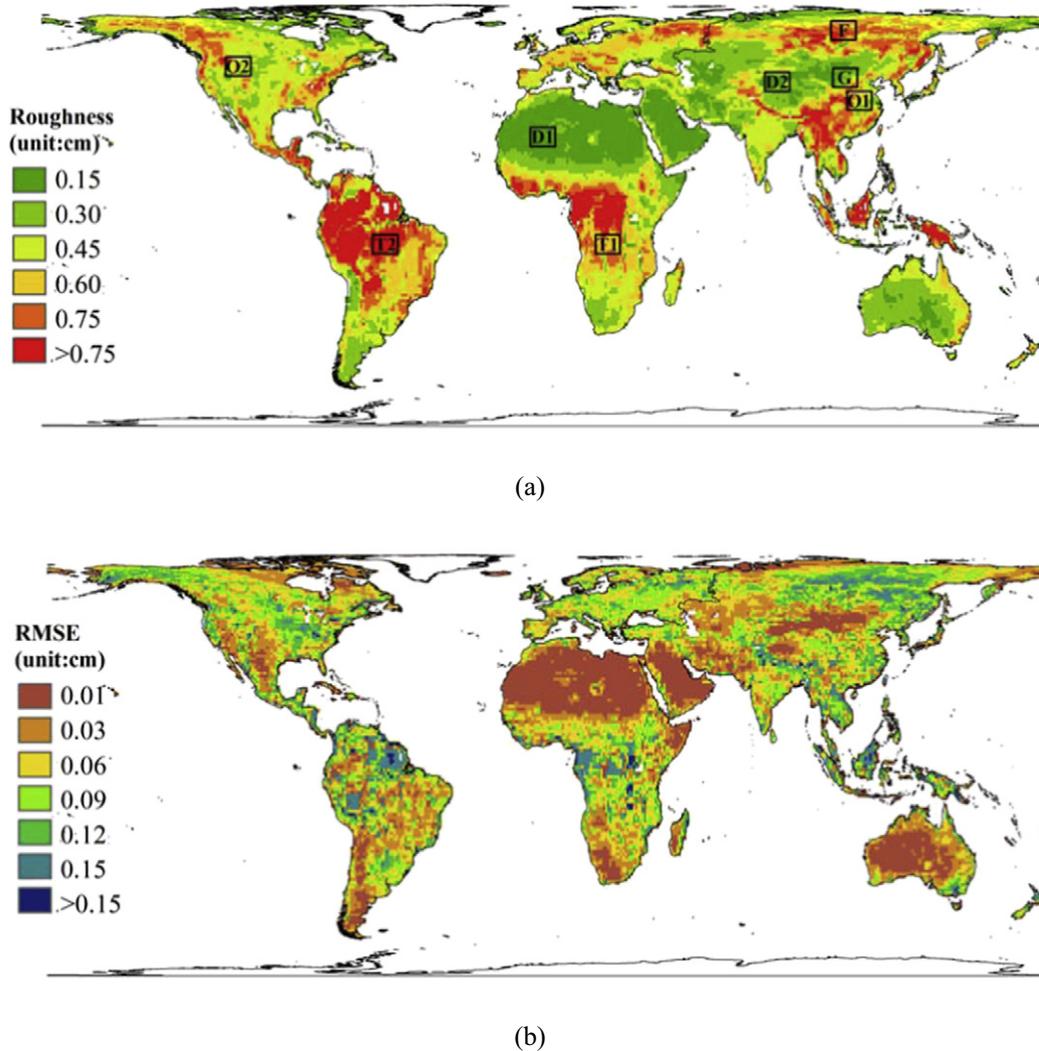


Fig. 3. Global annual average surface roughness and RMSE maps in 2009 (horizontal resolution: 25 km).

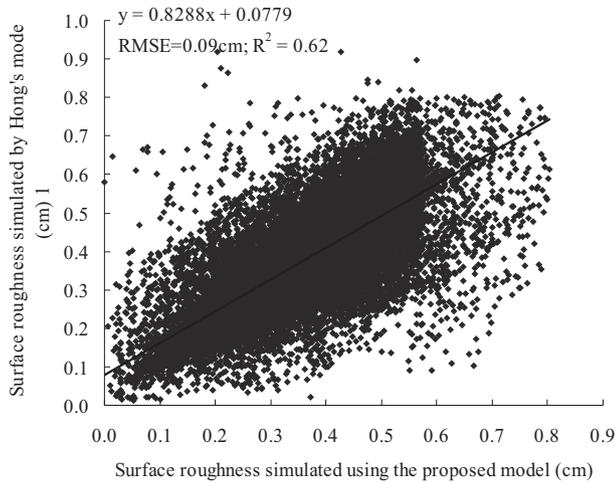


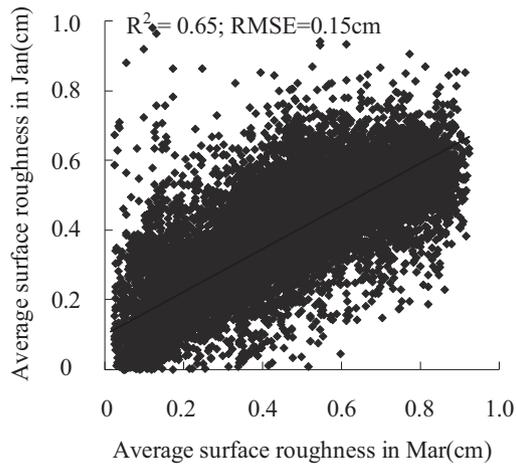
Fig. 4. Scatter diagrams between model-derived global monthly average surface roughness and those from Hong's (2010a) model in April, 2009.

Finally, the surface roughness retrieval models for AMSR-E C-band (Eq. (10-1)) and X-band (Eq. (10-2)) are established by substituting Eqs. (9-1), (9-2), (9-3), (9-4) into (8), respectively.

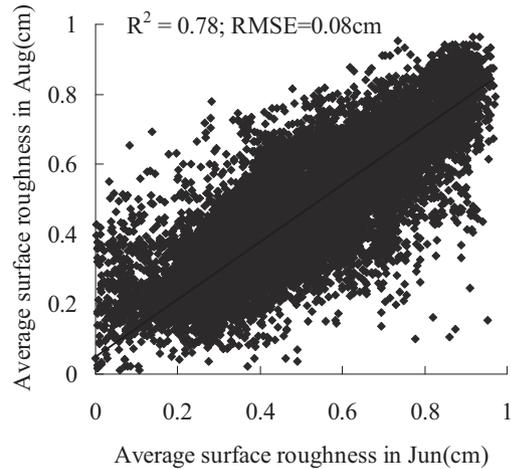
$$\begin{aligned} & (\text{MPDI}_{6.9} - 1 + 2Q) \times (0.7528 \times m_v + 0.0314) \\ & + (\text{MPDI}_{6.9} + 1 - 2Q) \times (0.7757 \times m_v^{0.4481}) \\ & = 2(\text{MPDI}_{6.9})^\alpha \times e^{\beta+h} \end{aligned} \tag{10-1}$$

$$\begin{aligned} & (\text{MPDI}_{10.7} - 1 + 2Q) \times (0.7117 \times m_v + 0.0284) \\ & + (\text{MPDI}_{10.7} + 1 - 2Q) \times (0.7619 \times m_v^{0.4610}) \\ & = 2(\text{MPDI}_{10.7})^\alpha \times e^{\beta+h} \end{aligned} \tag{10-2}$$

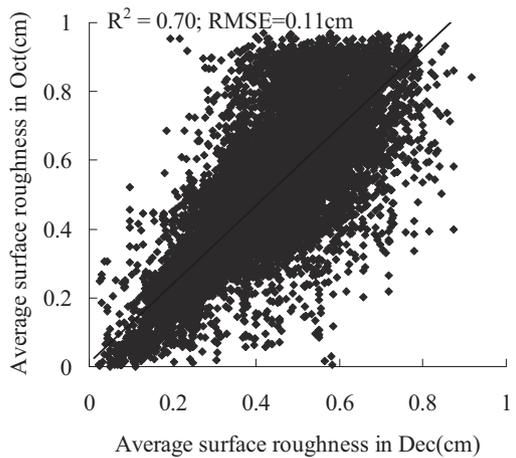
Because Eqs. (10-1) and (10-2) are two equations that contain two same variables (h and m_v), the surface roughness parameter (h) can be retrieved from AMSR-E C- and X-band MPDI after deleting the m_v variable by combining the two equations. Then, surface roughness (σ)



(a)



(b)



(c)

Fig. 5. Comparisons of the global monthly averaged surface roughness between January and March, June and August, and October and December in 2009.

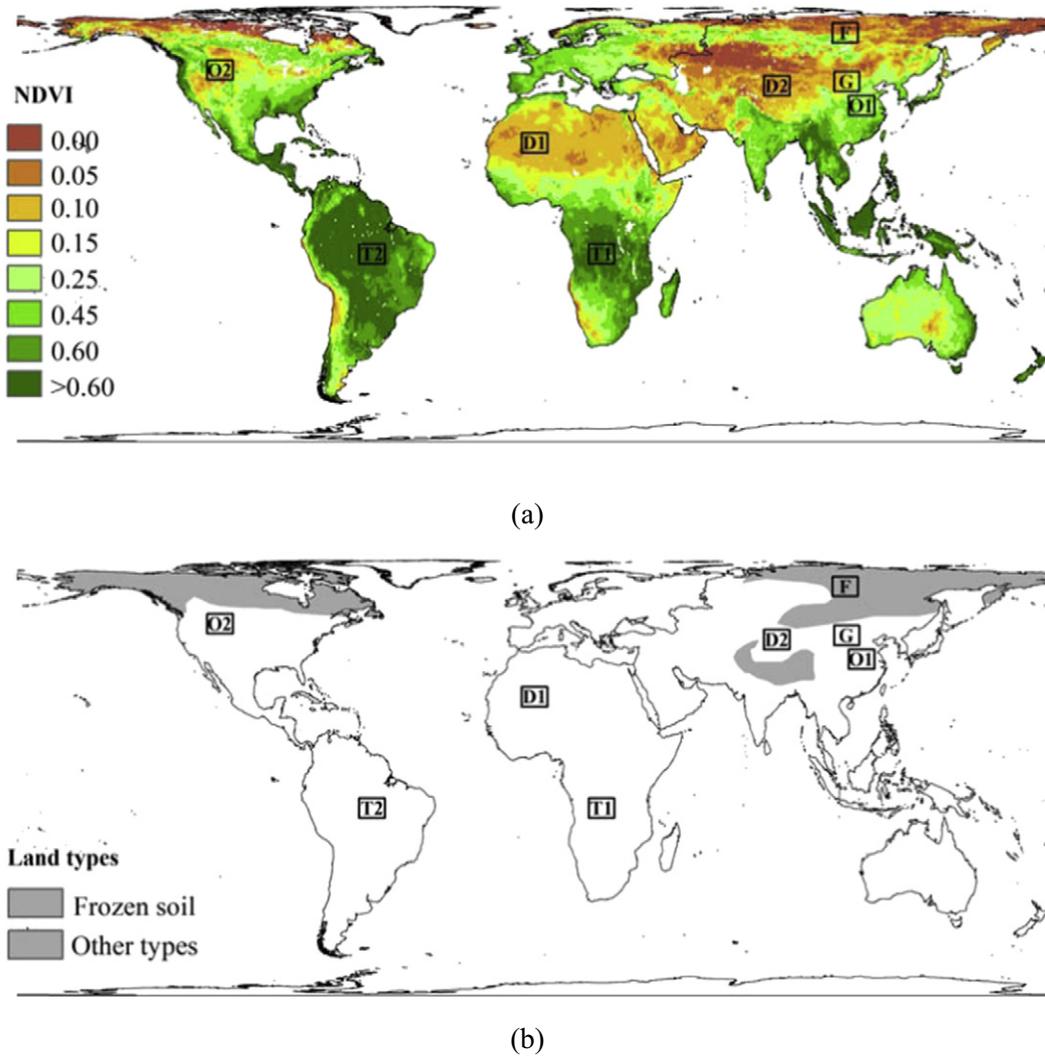


Fig. 6. The global annual mean Normalized Difference Vegetation Index (NDVI) map (horizontal resolution: 0.5 s) and freezing ground distribution map in 2009.

can be derived from h using Eq. (1). Values of the model parameters in Eqs. (10-1) and (10-2) are listed in Table 2.

4. Results and discussion

4.1. Global surface roughness mapping results and validations

The global annual average surface roughness map in 2009 is shown in Fig. 3a. The spatial horizontal resolution is 25 km as the same as that of AMSR-E T_p data. There are clearly visual correspondences that the Northern Eurasian continent, Northern Asia, Southwestern China, Central Africa, the western areas of North America, the northern areas of South America and the Eastern Australia are usually distributed with high surface roughness. While low surface roughness is mainly located at the Northern Africa, Central of Eurasian continent and Southern Australia. One-month global land surface roughness (224 168 samples in April, 2009), estimated using Hong's (2010a) method (Appendix A), is used to validate the retrieval accuracy of the proposed surface roughness model. Results show that the proposed model could effectively measure the surface roughness ranging from about 0.0 cm to 1.0 cm, which is similar as Hong's study. There is a strong linear relationship ($R^2 = 0.62$; $p < 0.001$) between them, with the root mean square error RMSE equals to 0.09 cm (Fig. 4). When the surface roughness of the input pixel is very high or very smooth, roughness values of neighborhood eight pixels are averaged and assigned to the input pixel. In

order to test the robustness of our global land surface roughness mapping method, a global RMSE map is also generated to show the spatial variations of the surface roughness retrieval errors (Fig. 3b). The retrieval RMSE ranges of <0.03 cm, 0.03–0.06 cm, 0.06–0.09 cm and >0.09 cm account for about 42%, 24.9%, 18.4%, 10.0% and 4.7% proportions of the global land areas, respectively. In other words, there is in total 85.3% of the global lands with the retrieval error smaller than 0.09 cm. There is one issue worthy mentioning here. When the vegetation cover becomes too thick for AMSR-E sensors to sense the roughness of the surface below, the retrieval model would not work well. This phenomenon has

Table 3
Land cover/land use and surface roughness conditions of the selected eight regions.

Number	Annual average NDVI	Land use types	Annual average surface roughness (cm)
D1	0.08	Desert	0.171
D2	0.03	Desert	0.188
T1	0.74	Tropical rainforest	0.664
T2	0.68	Tropical rainforest	0.867
F	0.11	Frozen soil with little vegetation cover	0.646
G	0.15	Grassland	0.292
O1	0.28	Deciduous vegetation cover	0.440
O2	0.22	Deciduous spare vegetation cover	0.472

been demonstrated in our recent studies for surface m_v and T_s simulations using AMSR-E T_b data (Chen et al., in press). A threshold value has been demonstrated that when the region's MPDI is smaller than 0.01, the vegetation cover is so dense that the microwave signals emitted by surface soils are impacted seriously when penetrated through a canopy. In other words, the AMSR-E sensors fail to sense the surface conditions effectively, which therefore leads to high retrieval errors.

4.2. Spatial and seasonal variations of the global surface roughness

We further compared the global surface roughness between different months in 2009. The scatter diagrams of global monthly averaged surface roughness between January and March, June and August, and October and December are presented in Fig. 5. Results show that the land surface roughness in March is generally higher than that in January, (Fig. 5a). Similarly, the land surface roughness in October is a bit higher than that in December (Fig. 5c). However, there is little difference between the land surface roughness in June and that in August (Fig. 5b). Results indirectly indicate that seasonal land surface cover changes might probably lead to time-series variations of the surface roughness conditions. In order to reveal the potential impact factors of global land surface roughness, several regions with different vegetation covers and land surface conditions are chosen for comparisons (Fig. 6 and Table 3).

Annual variations of the monthly average surface roughness in the eight regions are shown in Fig. 7. In general, the surface roughness of vegetated lands (T1, T2, O1, O2 and G) or freezing soil (F) is significantly higher than that of un-vegetated bare ground and desert (D1 and D2) (Table 3). The Sahara Desert (D1) in Northern Africa and Taklimakan Desert (D2) in Southern China are found to be covered by low surface roughness with slight seasonal variations in the whole year. The tropical rainforests in the Congo River Basin of Western Africa and Amazon Basin (T2) of Southern America are covered by high surface roughness with slight seasonal variations. The surface roughness of grassland in the Inner Mongolian Plateau (G) and freezing soil ground in Northern Asia (F) also varies little between different months. It is perhaps because that the land surface covers and land use conditions change slightly during the whole year. Contrarily, the deciduous vegetation regions in U.S.A. (O2) and Central China (O1) vary significantly. The reason might be that land surface vegetations become thicker in spring and summer, but defoliate in autumn and become snow-covered in winter. In summary, we can infer that vegetation cover and freezing soil will cause a rougher land surface, while snow cover will smooth the land surface.

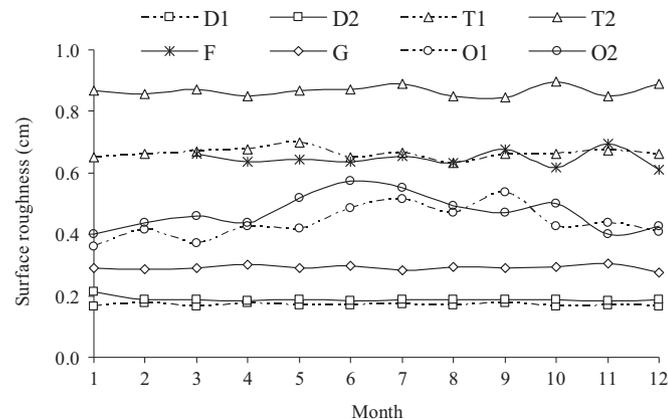


Fig. 7. Seasonal variations of monthly averaged surface roughness in 2009.

5. Conclusions

A global surface roughness simulation model is developed on the basis of the passive microwave radiative transfer equation using AMSR-E C- and X-band brightness temperatures. Validation results prove that the proposed surface roughness retrieval method is stable and accurate in characterizing the global land surface roughness. Analysis results indicate that vegetation cover, freezing soil and snow cover are important factors influencing the land surface roughness conditions. The surface roughness of vegetated lands or freezing grounds is significantly higher than that of un-vegetated bare grounds and deserts, while snow-cover regions always have lower surface roughness. What needs to be mentioned is that the proposed global surface roughness retrieval method only uses AMSR-E C- and X-band T_b without any ancillary data and might be fitted for global applications, but additional work using extensive in-situ data is still needed to validate the robustness of the surface roughness retrieval model.

Acknowledgments

We are grateful to the editor and anonymous reviewers for their valuable comments and suggestions, which have greatly improved the quality of this manuscript. This study is supported jointly by Jiangsu Key Laboratory of Agricultural Meteorology (grant no. KYQ 1207), the Guangdong Science and Technology Plan Fund (grant no. 2011B020313001) and Huizhou Science and Technology Plan Fund (2013B020006005).

Appendix A

Hong (2010a) ignored the atmospheric contributions on AMSR-E low frequencies (<10 GHz), took no consideration of volume scattering ($\omega \approx 0$), assumed that the surface soil temperature T_s equals to the vegetation canopy temperature T_c and then established the following rough reflectivity R_r equation.

$$R_r = e^{2\tau_c / \cos\theta} \left(1 - \frac{T_b}{T_s}\right) \quad (\text{a1})$$

Next, Hong (2010a) developed an approximate relationship between the V-polarized and H-polarized specular surface reflectivities with the same complex refractive index and derived the small-scale roughness model (Eq. (a2)) based on the W-C reflectivity model (Wang and Choudhury, 1981).

$$\sigma = \frac{\lambda}{4\pi \cos\theta} \sqrt{\ln \left(\frac{R_{r,V} \cos^2\theta}{R_{r,H}} \right)} \quad (\text{a2})$$

Following, the empirical linear relationship (Eq. (a3)) between the V-polarized T_b at 37 GHz $T_{b,37 \text{ GHz},V}$ and 1.25 cm T_s (De Jeu, 2003) was used to replace the T_s variable in Eq. (a1). So, small-scale roughness variable σ could be calculated from microwave T_b of AMSR-E low frequencies.

$$T_s = 0.861 \times T_{b,37 \text{ GHz},V} + 52.550 \quad (\text{a3})$$

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